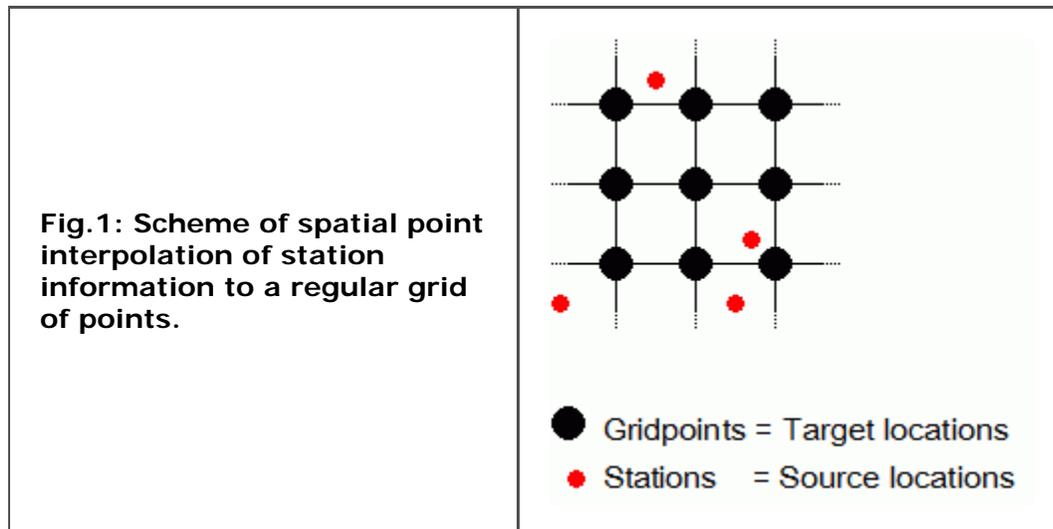


## 5.1. Introduction to Geostatistics and the spatial interpolation of weather and crop data.

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**The major aim of spatial interpolation is to get information of locations for which no observations are available.** The information is then interpolated to these locations from neighbouring locations for which the desired information is available. In agrometeorology one is usually interested in the spatial interpolation of meteorological data as well as crop yield data on a regular grid. The basic idea of spatial interpolation is depicted in Fig 1.



Usually the interpolated values at each grid point are regarded as representative for the surrounding square. However, one should keep in mind that the point estimate of the point interpolation is conceptually different from the area estimate of a block interpolation. In the latter case the target is not the value of a single location, but the area average of the square surrounding the target location. Most of the interpolation methods are designed to provide point estimates. Area averages can be obtained by averaging over a fine grid of point estimates.

Many questions are linked to the vast area of spatial interpolation of agroclimatic data. How many neighbouring stations should be used? How far may they be away to add information? How to mathematically obtain the estimate for the target location? Which methods are the best? How reliable are the results? The following sections aim to answer part of these questions and demonstrate how `New_LocClim`, the interpolation tool within the CM Box, solves the according problems.

### Which methods are available and what are their properties?

There is quite a variety of interpolation methods ranging from very simple and easy-to-use ones to very sophisticated strategies taking into account lots of supplementary information.

The variety of methods can be grouped with respect to certain features in order to easily get an idea of what they provide. The grouping criteria and their meanings are explained in Tab. 1.

Tab. 1: Grouping criteria for interpolation methods, short explanation and examples.

Criteria	Explanation	Example
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Scale free vs. scaled	A scale free method does not recognize the size of the system. Scale free methods do not distinguish the case where one station is 10m and the other 30m away from the case where one station is 100km and the other is 300km away. However, from an agroclimatical point of view there is a tremendous difference which should be taken into account. Scaled methods often use a scale parameter. If the chosen scale parameter is too large local variations are smoothed out. If it is too small, the information that may be contributed by stations further away is practically ignored.	Inverse distance interpolation is scale free. Shepards method is scaled.
Exact vs. smoothing	What happens if the target location falls exactly on the location of a station? An exact method would reproduce the observations. Smoothing methods take the observations of other locations into account and lead to a more averaged estimate.	Inverse distance interpolation is exact. Modified inverse distance interpolation is smoothing.
Parameter free vs. with parameters	If the user has not to provide any parameter to interpolate (and the method does not use built-in parameters) the method is called parameter free.	Nearest neighbour interpolation is parameter free.
Continuous vs. discontinuous	If for any 2 locations the interpolated values converge if the locations converge then the method is called continuous.	Nearest neighbour interpolation is discontinuous. Global polynomials are continuous.
Local vs. global	A method that takes all stations into account for each target location is called global. If it uses only part of the stations it is more or less local. Each method can be 'localized' by limiting the distance of stations to the target location or the number of stations to be used.	Nearest neighbour interpolation is the most local.
Extrapolating vs. non-extrapolating	A method is called non-extrapolating if the range of interpolated values does not exceed the range of observed values.	Kriging is extrapolating, inverse distance averaging not.

## Geostatistical Interpolation Methods

### Which method is the best?

There is no best method. You have to find the most adequate method with respect to the data given, station density and distribution, geographic conditions, season, degree of temporal integration (i.e. daily, decadal, monthly data) etc.

This is why the geostatistical methods which are available within the CM Box are described here in more detail.

### Nearest Neighbour

*Scale free, exact, parameter free, discontinuous, local, non-extrapolating*

This is by far the simplest method of spatial interpolation. For each grid point the value of the nearest station is taken.

### Inverse Distance Weighted Averaging (IDWA)

*Scale free, exact, 1 parameter, continuous with singular points, global, non-extrapolating*

IDWA is an averaging procedure. It gives each neighbouring station a weight which is proportional to a power of the inverse of the station distance. Thus, closer stations have more weight in the averaging procedure than stations that are further away. If the power exponent is large, closer stations get much more weight than stations that are further away. Therefore, with an increasing power exponent, IDWA converges to nearest neighbour interpolation. On the other hand, the smaller the exponent is, the further reaches the influence of each station. For exponents close to zero, the interpolated field becomes the area average with narrow spikes at station locations. Note that if a station is very close to the grid point it gets nearly all the weight. If a grid point and station are at exactly the same location IDWA is not defined because of division by zero. However, this is not a problem as in the very close vicinity of a station IDWA can be replaced by nearest neighbour interpolation.

### **Modified IDWA**

*Scale free, smoothing, 2 parameters, continuous, global, non-extrapolating*

One of the disadvantages of IDWA is that the influence of the closest station becomes rather high if this station is very close to the grid point compared to the other stations of the neighbourhood. This leads to the circus-tent like form of the interpolated field. In order to avoid this, Modified IDWA adds a displacement to each station distance. This means that even in the close vicinity of a station, this station does not get most of the weight, so the interpolated surface becomes smoother and the singular points (station locations) vanish. However, Modified IDWA is no longer an exact interpolation routine.

### **Cressmanns Method**

*Scaled, smoothing, 2 parameters, discontinuous, local, non-extrapolationg*

Cressmanns Method avoids some of the shortcomings of IDWA. Only stations within a circle around the grid point are used for the interpolation. Within this circle the weights of the stations depend on the distance from the grid point. A pre-given weight function, which stays high in the close vicinity of the grid point and becomes steeper with increasing distance and finally smoothly approaching zero, ensures that the weights are distributed with respect to a scale. This is a considerable advantage compared to IDWA. However, the radius of the circle has to be described, and the scale is fixed. If the radius is smaller than the closest station at a grid point the method gives no result. If the radius is large, the finer details of the climatological field may be smoothed out. Thus this method is not recommended where there are groups of stations with large gaps between them (a condition which is quite common in arid areas).

A modified version of Cressmanns Method allows the range to be broadened where the Cressmanns weights are high. This is done by an exponent. The higher it is, the higher is the weight of the stations further away compared to the closer ones.

### **Structure Function**

*Scaled, smoothing, n parameter, continuous, global, non-extrapolating*

The disadvantage of Cressmanns method is that stations outside the pre-given circle are not taken into account. Structure Functions avoid this disadvantage and simultaneously use the advantage of a pre-given scale. This means that if the neighbouring stations are all far away with respect to the scale they all get nearly the same weight. The point is that none of the far-away neighbours is considered to contain much more information about the local conditions than any of the others. However, the problem with this method is that the user has to provide a scale parameter.

Many monotonic functions can be used as Structure Functions. Within CM-Box two are provided: a Gaussian and an exponential. With the first, the influence of a station decreases over distance in a bell shape. The scale parameter gives the distance at which the power decreases to half its value. With the second structure function, the influence decreases exponentially. The scale parameter gives the distance at which the power of the station decreases to  $1/e$  (about 0.368) of its initial power of 1.

The advantage that Cressmann's method and Structure Functions have in common is that one knows exactly (or at least one can know) which part of the spatial high frequency variability is suppressed.

## Polynomials

*Scale free, smoothing, nonparametric, discontinuous, local, extrapolating*

This interpolation method is completely different from the previous ones. It fits a deterministic, two-dimensional function to the observations given by the neighbouring stations. The function value at the grid point represents the interpolated value. Since the function parameters are determined by a least-squares fit, it is by definition a smoothing method. Its smoothness depends on the number of neighbours which are used for the fit. It gives reasonable results with a high station density and in case of poor accuracy in the observed values. It has no scale property. Its major disadvantage is that if all the neighbouring stations are grouped to one side of the grid point the estimated value becomes a (non-linear) extrapolation of the local horizontal trend. Thus the estimates at the borders of a map may not be at all reliable.

## Shepards Method

*Scaled, exact, several parameters, discontinuous, local, non-extrapolating*

Shepards method is especially designed for climate interpolation. First a circle is drawn around the grid point containing, on average, seven stations. If there are less than four or more than ten stations within the circle the radius is changed so that the fifth nearest or the eleventh nearest station lies on the frame of the circle. This procedure ensures that the length scale is taken as a constant if possible. Only in regions with very high or very low station density the scale is changed to keep the number of stations used for the interpolation between four and ten. Within the circle a weight function is applied which monotonically decreases with the radius. In the next step the weights are altered with respect to directional isolation of the stations. This step is like a simple shadow correction. Since the weight function itself is like inverse distance weighted averaging in the close vicinity of a grid point, it would lead to singular points if a grid point and a station location coincide. Even if they do not coincide but lay close together, the grid point may become an unreasonable local maximum or minimum just because the closest station has an extraordinarily high weight. This situation is avoided by using a small local average-gradient correction. Finally, if one or more stations are very close to a grid point the simple average of these stations is taken as the interpolated value.

One advantage of Shepards Method is that it runs completely automatically. No parameters have to be set to obtain results. It automatically takes station density and grid size into account.

## Kriging

*Scales adapt automatically, smoothing, parameter free, continuous, global, extrapolating*

The basis of Kriging is a probabilistic concept with the following assumptions:

- The variable is random in space,
- Invariance of translation,
- Invariance of rotation.

Based on these assumptions a variogram is calculated. It gives the mean squared difference in the observed values as a function of the distance between the locations of observations. It can be shown that the estimate for the grid point can be expressed as a linear combination of the surrounding observations with weights following from the variogram. With the conditions given above the estimates are 'blue' (best linear unbiased estimators). This means that no other linear interpolation technique gives better results on average (smaller root mean square error). However, the above given assumptions mean that

- the variable should be stochastic only with no deterministic part in it,

- it makes no difference for the spatial correlation where you are,
- direction does not matter.

Only if these conditions are fulfilled, it be shown that simple Kriging leads to 'blue' results.

Apart from simple Kriging, many variations of the method exist. One of them is Co-Kriging where another variable is taken into account. This can be either another climatological variable or any other reasonable one, e.g. altitude. Satellite Enhanced Data Interpolation (SEDI) is a special application of Co-Kriging. It uses the spatial pattern of satellite images of cold cloud top duration for the interpolation of tropical precipitation fields.

### Thin-Plate Splines

*Scale free, exact, parameter free, continuous, global, extrapolating*

Thin-Plate Splines is a special kind of radial basis function. It can be seen as a deterministic method with a local stochastic component. The main condition is the minimum of curvature of the deterministic part. Another condition is that the interpolated surface has to meet the observations.

Also Thin-Plate Splines comes in a variety of versions. The deterministic part can be any arbitrary function. The so called smoothing Thin-Plate Splines allow a smoothing parameter to decide how strong the influence of the local stochastic part is. Note that minimization of curvature may lead to unreasonable extrapolations of local horizontal climate trends at the margins of the map.

## Deterministic structures

Aside the stochastic part of agroclimatic variables there is the deterministic fraction which can be rather developed. The fraction of spatial variance of a variable that can be attributed to deterministic structures varies from variable to variable and with respect to the degree of integration. In the extreme case of long-term averages of monthly mean temperatures the overwhelming majority of spatial variability within hundreds of kilometers may be explained by altitude alone. There may also be a clear horizontal gradient. Before interpolating agroclimatic data it should be analysed which fraction of variance is explained by either vertical or horizontal gradients. If they explain a considerable fraction of variance they should be treated separately as explained in the following subsections.

### Altitude Correction

Either a look on the observation-altitude graph or a look on the explained fraction of variance by altitude (or both) clarifies whether altitude matters or not. If it matters it should be taken into account by transforming the observations via linear regression to values that would have been observed at the station location if it had the same altitude as the target location. Then, in a second step the spatial interpolation of the pre-transformed observations may be performed. Aside an improvement of the interpolation this yields to an understanding of the effect on altitude on the variable under consideration. However, several problems may arise in data sparse areas with altitude regression. They are discussed in the section Warnings.

### Horizontal Gradients

Some of the interpolation methods take deterministic spatial structures into consideration (polynomials, thin-plate splines). The easiest possible spatial structure is a horizontal gradient (trend), and if the region which is covered by the neighbourhood of stations is small enough, even more complicated regional spatial structures may be well approximated by local planes. Therefore, we suggest local gradients to be taken into consideration in the same way as is done with the altitude correction.

However, we not recommend to apply horizontal and altitude regression in one step since both can be collinear and a strange and meaningless mixture may follow. If both altitude and horizontal gradients should be applied, we suggest dealing with them separately. This means, that first the altitude dependency should be approximated and then the horizontal gradients.

## How good is spatial interpolation?

The quality of spatial interpolation of agroclimatic data depend on the station density, station distribution and the spatial representativeness of the observed variable. The representativeness may be expressed in a scale length which is the distance of two stations which have half of their variability in common. Representativeness depends on the variable itself, the season, degree of temporal integration (daily, decadal, monthly, long term averages, etc.), the climatic regime, the geographic and topographic conditions (flat land, hills, mountains, continentality, maritimity, etc.). Some estimates of scale lengths exist in literature. However, usually this information is missing. Therefore one needs another way to measure the quality of interpolation.

A simple and conservative way measuring the quality of spatial interpolation is the jackknife error.

Instead of interpolating the climate at grid points we interpolate it to the station locations and leave the observation at the station out for the interpolation at its location. Then the error of interpolation at each station is just the difference of the interpolated and the observed value. It can be seen as a property of the station with respect to the method used and the neighbouring stations. Averaging all these station-based errors leads to an estimate of the interpolation bias. The sum of squared deviations from this bias gives an average station-based error.

However, imagine the case where all stations but one are close together and their observations are rather similar. One station however, is far away and has a completely different value. In this case all but one will have a small Jackknife error and one will have a large one. Should we simply average the errors? If we are interested in area specific answers we should not. All the small Jackknife errors together represent a small sub-region, whereas the separated station may represent the major part of the total area. Thus the averaging should include the region for which the corresponding Jackknife error is representative. We therefore suggest to attribute a jackknife error to each gridpoint by nearest neighbour interpolation and to average afterwards over all gridpoints to get an area averaged Jackknife error as well as an area averaged standard deviation. In this way the Jackknife error also becomes a function of the grid size. Thus, one may wish to look at both the station specific (grid size independent) root mean square Jackknife error as well as the area specific one.

## How many neighbouring stations should be used?

The stations around a gridpoint are used to obtain an interpolated value at the gridpoint. The basic question is: How many stations should be used for the interpolation? The more stations that are used, the more likely the effect of a local outlier will be smoothed out. But does more data really mean better results? Using more stations for the interpolation means more data are used, but they are from stations that are further away. Do they really add information? Do they matter at all? Or do they worsen the results because they add information from a distant location? Some methods give the stations that are further away much less weight in the interpolation procedure thus these stations make hardly any difference to the results so using them is optional and little more than a waste of time. Other methods exist which automatically fix the neighbourhood of stations to be taken into account. In all other cases the user can decide how many stations to use. We suggest to keep the number of neighbouring stations small. The default value of `New_LocClim` is 10. This sounds low, but one should bear in mind that the aim is to estimate the climate at a local grid point and that this neighbourhood is generated for each grid point. In some cases the tenth nearest station can be rather far away and add information of another climate regime. It may be best to look at the spatial distribution of the stations and decide from there how many neighbours are optimal. The Jackknife error can be a helpful tool with respect to this decision. However, `New_LocClim` can deal with up to 100 neighbours.

There is also the possibility to limit the number of neighbouring stations by distance. Note that limiting the distance too strictly, grid points may occur where no stations fulfil the conditions of the neighbourhood. This will lead to blank areas on the interpolated climate map.

## Shadowing

Seen from the target location some stations may be hidden behind others. They may not add new information of the climate at the target location. Fig. 2 illustrates the effect of such groups of stations. Pure distance weighting methods are prone to these situations. Therefore in case of large grouping of stations one should either eliminate some of them to obtain a more homogeneous distribution or one may apply a shadowing that gives less weight to the stations which are in the shadow of others. Kriging methods automatically take care of this effect and can not be improved by shadowing. This is one of the strengths of kriging.

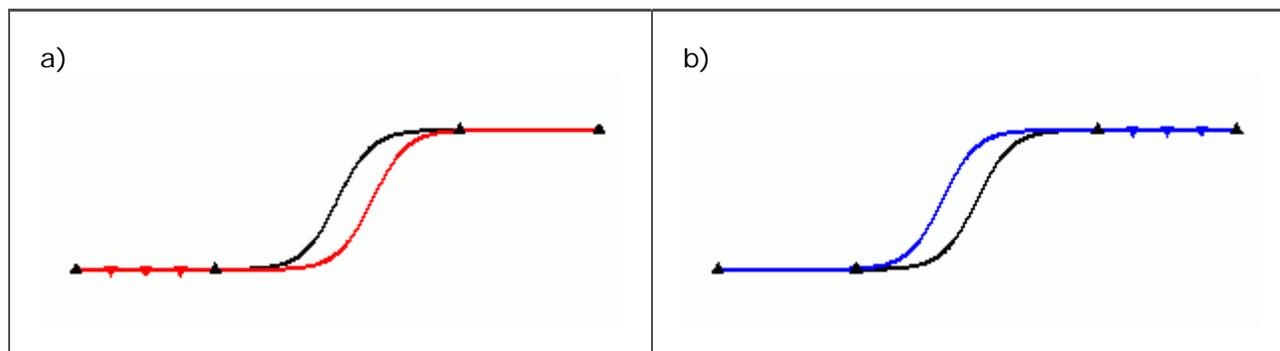


Fig. 2. Effect of stations behind other stations on the interpolation surface. Black lines are interpolation results if only 4 stations (black triangles) are used. In panel a) three more stations are added on the left side, in panel b) three more stations are added on the right side. These stations do not add information, but alter the interpolation results as sketched by the coloured lines.

## How to detect strange data?

If one assumes that the observed values for each month are the sum of a horizontal function, a vertical function, and random noise one can detect outliers. As a first step all data may be reduced to the same altitude and location by the functions fitted to the data. In the next step the value furthest away from the average is assumed to be an outlier and thus extracted to obtain unbiased average and variance of the residual ensemble. The latter can be used to calculate thresholds which are exceeded very rarely by chance. If the probability to exceed a threshold in one trial is given as  $q$  and one assumes independent trials (i.e. independent observations at the neighbouring stations) then one can apply a Poisson model to calculate the probability that one in  $n$  trials ( $n$  = number of neighbouring stations) is as far away from the observed average as the assumed outlier is. If this probability is less than a certain threshold (10% say) the observation should be regarded as strange. In the single point mode of New\_LocClim the station, month and variable of the strange observation is provided.

In the map mode of New\_LocClim each station can be a neighbour for many grid points. It may provide strange data with respect to some of these cases. Thus, strangeness of a station is provided in this case as the ratio of cases where the station seemed to provide strange data. The results are provided in the sub menu manually of the menu item stations and as station information on the map.

## A practical guide

As a general strategy for the interpolation of agroclimatic data we recommend to first investigate the data in order to answer the following questions:

1. Do I have enough data?

This can be recognized by an acceptable Jackknife error.

2. Is the station distribution acceptable?

Draw a map of the station locations. If you see large gaps come to think of it. Do they

matter? Is there any peculiarity with the gaps (e.g. a mountain range)? In this case one should not simply interpolate over the gap but try to close it by obtaining data from the area affected. Do not forget to care for the marginal areas. Try to obtain and take into account stations surrounding your area of interest in order to get reliable results also at the margins of the map.

### 3. Are there strange data?

Draw the observations in a map either as values or as coloured points. Outliers become visible this way. But be careful. Sometimes the only mountain station on a map provide valuable information since the observation is quite different from the observation at all other stations. This does not mean that it is an outlier. Use the strange data warning of New\_LocClim. Drop obviously wrong values.

### 4. Do the data depend on altitude?

Draw a map of explained variances by altitude regression and a map of regression coefficients (slopes). If the explained variance is high and the regression coefficients are meaningful then apply altitude regression before applying geostatistical methods. The maps of explained variances and regression coefficients are valuable agroclimatic information. Experiment with the number of stations used to calculate local correlation and regression coefficients.

### 5. Do the data show a horizontal gradient?

This can be checked in the same way as the altitude dependency. If it is important it should be taken into consideration. Clarify which geostatistical interpolation method to prefer and find the optimal parameter values.

### 6. Do the stations cover the whole range of altitude of the grid points?

The altitude range covered by the stations should cover the range of altitude of the grid points. Otherwise the interpolation is in fact an altitude extrapolation. An altitude observation plot helps to see whether this is a problem. Truncated altitude regression may avoid altitude extrapolation. Maps of local altitude extrapolation may be drawn. Clarify which geostatistical interpolation method to prefer and find the optimal parameter values.

### 7. Is any of the grid points lonely?

If the group of neighbouring stations does not surround the target location but falls at one side of it, the target location may be called lonely and the interpolation to it may be seen as a horizontal extrapolation. A map of horizontal extrapolation warning may be drawn to clarify this.

### 8. Clarify which geostatistical interpolation method to prefer and find the optimal parameter values.

Some methods are based on features which are less adequate with respect to some variables. As an example thin-plate splines minimize curvature. The spatial field of precipitation however may have regions with high curvature and steep gradients. Large precipitation and no precipitation may be in very close vicinity. Therefore one might expect that thin-plate splines is not the optimal method for the interpolation of precipitation. However, superior to the theoretical investigations of this kind is trial and error. The method with the least Jackknife error and reasonable interpolated maps should be used. In order to estimate Jackknife errors, it is not necessary to do all the interpolations with the final resolution. Therefore this investigation may be less time consuming as one expects. Note that Jackknife errors of smoothing methods are usually lower compared to exact methods. Experiment with the number of stations used to calculate local correlation and regression coefficients.

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